PALYNOLOGY OF THE MESOZOIC SUCCESSION OF THE KALA CHITTA RANGE PAKISTAN

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ABSTRACT:-A total of one hundred six samples of the Mesozoic succession of the Kala Chitta Range were processed for the palynological analysis out of which only twenty one samples proved to be productive. Palynoflora recovered from the productive samples comprise thirty one form genera belonging to the thrity three form species. Eighteen triletes, two monoletes, eight bisaccates, two acritarchs and one monosaccates constitute eighteen genera. Palynoflora is less diverse exhibiting little compositional change. Lundbladispora willmotti and L. microconata are age diagnostic for the Upper Triassic Kingriali Formation. Aequitriradites triangulatus is typical of the Upper Jurassic-Lower Cretaceous Chichali Formation. Palynoflora are abundant in the Lower Cretaceous Lumshiwal Formation containing Succinctisporites grandior, Gordodinium alberti, Leptodineum eumorphum, Faveosporites subtriangularis, Densoisporites nejburji, Enzonalasporites vigens, Minutosaccus crenulatus, Ovalipollis ovalis, Podocarpites ellipticus and Aequitriradites triangulates. The Upper Cretaceous Kawagarh Formation contains Ptrospermella australliensis of Campanian age. The present contribution is the first comprehensive study of its kind.

INTRODUCTION

The Kala Chitta Range is a part of the foreland-fold and thrust belt which forms the northern border of the adjoining hydrocarbon bearing Potwar Basin (Fig. 1). The sampled localities have been shown in Fig. 2. The abstract of this paper forms part of the abstract volume of the 33rd International Geological Congress, Oslo, Norway. (Butt et al. 2008).

The Mesozoic succession of the Kala Chitta Range as formalized by the Stratigraphic Committee of Pakistan (Fatmi, 1973) has been tabulated here as follows.

| | Nomenclature of the Stratigraphic Committee of Pakistan (Fatmi 1977) | | | | | | | | | |
|-------------------------------------|---|--------------------------------|--|--|--|--|--|--|--|--|
| Upper Paleocene | Lockhart Limestone (Thanetian) Sequence Boundary (Absence of Maastrichtian-Danian) | Hill Limestone | | | | | | | | |
| Upper Cretaceous | Kawagarh Formation (Coniacian to Campanian) Sequence Boundary (Absence of Cenomanian-Turonian) | Kawagarh shales | | | | | | | | |
| Lower Cretaceous | Lumshiwal Formation | Giumal sandstone | | | | | | | | |
| Upper Jurassic- Lower Cretaceous | Chichali Formation Sequence Boundary (Lateritic Crust) | Spiti shale | | | | | | | | |
| Middle Jurassic | Samana Suk Limestone | Kioti Limestone | | | | | | | | |
| Lower Jurassic | Datta Formation | Ferruginous beds in the kiotos | | | | | | | | |
| Triassic | Kingriali Formation Chak Jabbi Limestone Mianwali Formation | Kioto Limestone | | | | | | | | |

Base not exposed by virtue of the Thrust Fault



Fig. 1. Location map of the Kala Chitta and Samana Ranges (after Davies and Pinfold,1937)



Fig. 2. Location map of the study area of Kala Chitta Range.

Palynomorphs are preserved in sediments under great variety of environments. During sedimentation they behave as very small particles like fine sand and silt. Coarse sandstones, oxidized and weathered sediments, dolomites and sometimes even limestones are poor containers of these microfossils. But some workers have reported well preserved palynomorphs in carbonate environments (Scott et al., 1985b; Masood et al., 1995, Blome and Albert, 1985).

Palynomorphs recovered during the present investigation are poorly preserved. In most cases it was difficult to apply correct identification beyond generic level. However, only those palynomorphs are included in this dissertation that were to some extant moderately preserved. Systematic section includes technical description of only selected palynomorphs most of which also appear in table No. 1.

A total of one hundred and six samples of the Mesozoic succession of the Kala Chitta Range were processed for palynological analysis. Only twenty one samples proved to be productive. The details are shown in table No. 1. It is quite evident from this table that the most productive samples belong to the Lumshiwal Formation, whereas, the Samana Suk Limestone and the Kingriali Formation were least productive.

METHODOLOGY

Special strew mount preparations were made (sandwich mounts between two cover slips) to allow morphographic examination of palynomorph from two angles i.e., proximal and distal. Descriptive features of each palynomorph species encompassing morphographic characters appearing in the systematic palynology were registered as they appeared in the actual specimen under oil immersion objective upon careful L-O analysis. Such features, however, may not be properly depicted in the photomicrographs provided due to poor state of preservation (e.g. crumpling, variously folding, differential decay).

REMARKS

Very few specimens allowed identification up to genus or species level. All palynomorphs posing problems regarding correct identification due to poor preservation were excluded from the study.

Palynoflora recovered from the productive samples comprise thirty one form genera belonging to thirty three form species. Of the thirty one genera 18 belong to triletes 2 to monoletes 8 to bisaccates 2 to acritarchs and one to monosaccates. Palynoflora is less diverse exhibiting little compositional change across the respective formation. Most miospores were of restricted origin, occurring only in one or two formations and very few long ranging types were encountered.

Consideration of Table No. 1 reveals that *Leiotriletes* and *Laevigatotriletes* are the only two long ranging genera occurring almost in all samples, whereas, *Alisporites grandis, Pinuspollenites sp., Affropollis jordinus, Inaperturopollenites sp.* and *Pterospermella australiensis* were the next long ranging types commonly shared by two or more formations. *Punctatisporites* which otherwise is a long ranging form genus (Devonian to Neogene) was found to be restricted only to the basal part of the Lumshiwal Formation.

PALYNOSTRATIGRAPHY

Kingriali Formation (Upper Triassic)

The Kingriali Formation contains four palynospecies of restricted origin, viz., *Lundbladispora willmotti*, *L. microconata*, *Goubinispora sp.* and *Indotriletes korbaensis*. The first three are age diagnostic, mostly restricted to later part of the Middle Triassic (Jansonius and Hills, 1976).

Samana Suk Limestone (Middle Jurassic)

No age diagnostic or the palynomorph of restricted origin was found in this formation. Apart from the general long ranging types *Leiotriletes* and *Laevigatosporites* only few *Mycorrhizae* were found here.

Chichali Formation (Upper Jurassic-Lower Cretaceous)

This formation is slightly more productive than the Kingriali and the Samana Suk Formations. It contains *Leptolepidites eparcornatus* as the restricted form. *Impardecispora cf. parvelentus* (Plate No. 1c) may be another such type, but extending to the lower most part of the overlying Lumshiwal Formation. Age diagnostic palynomorphs for this formation are *Aequitriradites triangulatus* and *Frangospora sp.* that are peculiar for Early Cretaceous or some time Latest Jurassic (Jansonius and Hills, 1976).

Lumshiwal Formation (Lower Cretaceous)

The Lumshiwal Formation is palynologically the richest horizon investigated during the present study. Palynoflora recovered from this formation is most abundant and diverse as compared to other formations. Punctatisporites sp. (Plate No. 4e), Succinctisporites grandior (Plate No. 4a), Gordodinium alberti (Plate No. 4c), *Leptodineum eumorphum* 2e), (Plate No. Foveosporites subtriangularis (Plate No. 2f), Densoisporites nejburji (Plate No. 2a), Enzonalasporites vigens (Plate No. 4d), Minutosaccus crenulatus (Plate No.

 Table 1

 The frequency of occurrence of the palynomorphs in the Mesozoic sucession of the Kala Chitta Range.

| | | TRIASSIC | JURA | SSIC | JU | JURASSIC/ CRETACEOUS | | | | | | | | | | | | | | | | |
|------------------|--------------------------------|----------|-------|------|-----|---|----|------------|-------------|-----|----------|------|----------|-----|--------------|------|----------|-----|------|------|-----|-----|
| | | | | | CF | RET | | EO | | | | | | | | | | | | | | |
| | AGE | WINCOLA | G A M | | | | 5 | T T | | | | | 1.01 | | 7 A T | | | | 17 | | | |
| FORMATION LI SUK | | | | | | | | | AL KAWAGARH | | | | | | | | | | | | | |
| | Sample No | 576 | 703 | 1029 | 65 | 66 75 18 14 14 200 210 661 66 66 75 197 | | | | | | | | 188 | 1236 | 1237 | 170 | 171 | | | | |
| S.No | PALYNOMORPH | 570 | 705 | 1027 | 8 | 9 | 4 | 1 | 8 | 9 | 207 | 210 | 001 | 5 | 6 | 6 | 107 | 100 | 1230 | 1237 | 170 | 1/1 |
| 1 | Affropollis jordinus | | | | | | | | | | | | 13. 5 | 2 | | 28 | | 10 | 4 | 3.5 | | |
| 2 | Densoisporites nejburgii | | | | | | | | | | 10. 5 | 15 | 6.5 | 1 | 13 | 4.5 | | | | | | |
| 3 | Alisporites grandis | | | | | | | 8 | 5 | 8 | 4 | 10 | 3 | 3.5 | 9 | 19 | 21. 5 | 4.5 | 8 | 1 | 2.5 | 2 |
| 4 | Alisporites bilateralis | | | | | | | 2 | 18 | 1.5 | 6.5 | 10.5 | 3.5 | 19 | 2.5 | 2.5 | 5 | 31 | 2 | 3.5 | 1 | 1 |
| 5 | Enzonalasporites vigens | | | | | | | | 3 | 1.5 | 11 | | | | | | 4.5 | | | | | |
| 6 | Frangospora sp. | | | | | | | *5 1 | 1 | 1.5 | 1.5 | 9.5 | 2 | 16 | 10 | | 21 | | | | | |
| 7 | Inaperturopollenites sp. | | | | | | | | | | | | | | 1.5 | 2.5 | | 3.5 | 2 | 1.5 | 1.3 | |
| 8 | Leptolepidites eparcornatus | | | | 4.5 | 2.5 | 6 | 10 | | | | | | | | | | | | | | |
| 9 | Leiotriletes sp. | | 85 | 90 | 69 | 43 | 29 | 10 | 10 | 17 | 2.5 | 3.5 | 9.5 | 2 | 8.5 | 1.5 | 6 | 9 | 2.5 | 2 | 4 | 19 |
| 10 | Laevigatosporites sp. | | 10 | 8 | 18 | 39 | 51 | 8 | 3 | 2 | 7 | 8.5 | 17 | 19 | 1 | 16 | 4 | 4.5 | 1 | 1 | 3 | 20 |
| 11 | Minutosaccus crenulatus | | | | | | | | 2 | 3 | 4 | 4 | 8.5 | | | | 5 | | | | | |
| 12 | Ovalipollis ovalis | | | | | | | | 16 | 4.5 | 3.5 | | | | | | 4 | | | | | |
| 13 | Podocarpites ellipticus | | | | | | | | 8 | 5.5 | 1 | 2 | 7 | 16 | 8 | 2.5 | 1.5 | 19 | | | | |
| 14 | Punctatisporites sp. | | | | | | | | | 28 | 11. 5 | 6.5 | 2 | | | | 1.5 | | | | | |
| 15 | Succinctisporites grandior | | | | | | | | 12 | 2 | 9 | | | | | | 4.5 | | | | | |
| 16 | Valiasaccites validus | | | | | | | | | | | | | 2 | 19 | 55 | | 1.5 | | | | |
| 17 | Leptodineum eumorphum | | | | | | | | 2 | | 13 | | | | | | 9.5 | | | | | |

Table No. 1:Continued from pre-page.

| S.No | Sample №. PALYNOMORPH□ | 576 | 703 | 1029 | 65 8 | 66 9 | 75 4 | 181 | 14 8 | 14 9 | 20 9 | 21 0 | 66 1 | 66 5 | 66 6 | 756 | 18 7 | 88 | 123 6 | 123 7 | 170 | 171 |
|------|-----------------------------------|-----|-----|------|---------|---------|---------|-----|---------|---------|---------|---------|---------|---------|---------|------|---------|-----|----------|----------|-----|-----|
| 18 | Gordodinium alberti | | | | | | | | 9 | 2.5 | 2.5 | | | | | | 2.5 | | | | | |
| 19 | Foveosporites subtriangularis | | | | | | | | | | | 18 | 8.5 | 1.5 | 1 | 11.5 | | | | | | |
| 20 | Crybelosporites pannuceus | | | | | 7 | 4 | 3.5 | 1 | 3.5 | | | | | | | 3 | | | | | |
| 21 | Verrucosisporites narmianus | | | | | | | | | | | | | | | | | | 34 | 29 | | |
| 22 | Verrucosisporites densus | | | | | | | | | | | | | | | | | | 20 | 14 | | |
| 23 | Cyclobaculisporites minutus | | | | | | | | | | | | | | | | | | 15 | 23 | 56 | 49 |
| 24 | Inferopollenites sp. | | | | | | | | | | | | | | | | | | 5 | 4 | | |
| 25 | Densipollenites densus | | | | | | | | | | | | | | | | | | 1.5 | 2 | 10 | 5 |
| 26 | Pinuspollenites sp. | | | | | | | | 1.5 | 1 | 1.5 | 3 | 10 | 1.5 | 19 | 1.5 | 1 | 10 | 2 | 2.5 | 1 | 3 |
| 27 | Aequitriradites triangulatus | | | | | 4 | 5 | 2 | 1 | 2 | | | | | | | 1 | | | | | |
| 28 | Lundbladispora willmotti | 53 | | | | | | | | | | | | | | | | | | | | |
| 29 | Lundbladispora microconata | 2 | | | | | | | | | | | | | | | | | | | | |
| 30 | Indotriletes korbaensis | 34 | | | | | | | | | | | | | | | | | | | | |
| 31 | Impardecispora cf. parvelentus | | | | 3.5 | 1 | 1.5 | 2.5 | 1.5 | | | | | | | | | | | | | |
| 32 | Pterospermella australiensis | | | | | | | | | | | | | | | | | | 3 | 12 | 9 | |
| 33 | Goubinispora sp. | 9.5 | | | | | | | | | | | | | | | | | | | | |
| 34 | Mycorrhiza | | 4.5 | 2 | 4.5 | 1.5 | 1 | 1 | 2 | 4 | 6 | 7.5 | 1.5 | 13 | 4.5 | | 1.5 | | | | | |
| 35 | Acritarchs | | | | | | | 1 | 4 | 3 | 1.5 | 1 | 1 | 1.5 | 1 | 3 | 1 | 5 | | | | |
| 36 | Dinoflagellate | | | | 1 | 2 | 1 | | | | | 1 | 2 | 1 | 2 | | 1 | 0.5 | Ī | | | |
| 37 | Algal remnants | | | | | | | | | 1.5 | 2 | 1 | 2 | 1 | 1 | 1 | 1.5 | 0.5 | | | | |

PLATE 1



- Fig. a Aequitriradites triangulates, Lumshiwal Formation. KQ. 188, Dheri Kot
- Fig. b Laevitriletes sp., Chichali Formation. KQ. 669, Burjianwala Laman
- Fig. c Impardecispora cf. parvelentus, Chichali Formation. KQ. 754, Chapra
- Fig. d Bisaccate, Lumshiwal Formation. KQ. 188, Dherikot
- Fig. e Indotriletes korbaensis, Chichali Formation. KQ. 754, Chapra
- Fig. f Pterospermella australiensis, Kawagarh Formation. KQ. 170, Choi
- Fig. g Alisporites grandis, Chichali Formation. KQ. 669, Burjinwala Laman
- Fig. h Podocarpites sp., Chichali Formation. KQ. 754, Chapra

PLATE 2



- Fig. a Densoisporites nejburji, Lumshiwal Formation. KQ. 665, Burjianwala Laman
- Fig. b Leiotriletes sp., Lumshiwal Formation. KQ. 208, Daurdad
- Fig. c Leptolepidites eparcornatus, Lumshiwal Formation. KQ. 187, Dherikot
- Fig. d Alisporites grandis, Chichali Formation. KQ. 181, 1/2, F 2/7, Bata
- Fig. e Leptodineum eumorphum, Lumshiwal Formation. KQ. 148, 6/3 F 3/24a, Togowala
- Fig. f Foveosporites subtriangularis, Lumshiwal. KQ. 208, 1/1, F 1/35, Daurdad
- Fig. g Minutosaccus crenulatus, Chichali Formation. KQ. 669, 4/1, F 3/2a, Burjianwala Laman
- Fig. h Alisporites bilateralis, Chichali Formation. KQ. 181, 4/3, F 2/8, Bata
- Fig. i Valiasaccites validus, Lumshiwal Formation. KQ. 147, 1/1, F 3/14, Togowala
- Fig. j Laevigatosporites sp., Lumshiwal Formation. KQ. 148, 3/3, F 3/20a, Togowala
- Fig. k Laevigatosporites sp., Lumshiwal Formation. KQ. 148, 6/3, F 3/23a, Togowala

PLATE 3



Lumshiwal Formation

- Fig. a Succinctisporites grandior. KQ. 149, Togowala
- Fig. b Inaperturopollenites sp. KQ. 148, Togowala
- Fig. c Ovalipolis ovalis. KQ. 149, Togowala
- Fig. d Crybelosporites sp. KQ. 209, Daurdad
- Fig. e Crybelosporites pannuceus. KQ. 149, Togowala
- Fig. f Afropollis jardinus. KQ. 756, Chapra

PLATE 4

Lumshiwal Formation

- Fig. a Podocarpites ellipticus. KQ. 148, Togowala
- Fig. b Palambages sp. KQ. 149, Togowala
- Fig. c Gordodinium alberti. KQ. 149, Togowala
- Fig. d Enzonalasporites vigens. KQ. 148, Togowala
- Fig. e Punctatisporites sp. KQ. 149, Togowala
- Fig. f Frangospora sp. KQ. 148, Togowala
- Fig. g Endogone sp. KQ. 149, Togowala

PLATE 5

Kawargarh Formation

- Fig. a Cyclobaculisporites minutes. KQ. 170, Choi
- Fig. b Verrucosisporites narmianus. KQ. 1236, Surg
- Fig. c Pinuspollenites sp. KQ.1225, Dherikot
- Fig. d Verrucosisporites densus. KQ. 170, Choi
- Fig. e Palambages sp. (Acritarch). KQ. 1237, Surg
- Fig. f Leiotriletes sp. KQ. 170, Choi
- Fig. g Densipollenites densus. KQ. 1237, Surg
- Fig. h Infernopollenites sp. Kq. 170, Choi
- Fig. i Sclerocystis sp. KQ. 158, Daurdad

2g), *Ovalipollis ovalis* (Plate No. 3c) and *Podocarpites ellipticus* (Plate No. 4a) are of restricted origin, whereas, *Frangospora sp.* (Plate No. 4f) and *Aequitriradites triangulatus* (Plate No. 1a) are age diagnostic.

Kawagarh Formation (Late Cretaceous)

Kawagarh Formation was poorly productive. Apart from Leiotriletes and Laevigatotriletes that are the general long ranging types for all the formations only four palynomorphs species are encountered here. These are also commonly shared by the Lumshiwal Formation, viz., Affropollis jordinus (Plate No. 3f), Alisporites grandis, Inaperturopollenites sp. (Plate No. 3b) and Pinuspollenites sp. (Plate No. 5c). Apart from these commonly shared types, the Kawagarh Formation also contains Ptrospermella australiensis (Plate No. 1f) that is age diagnostic for Campanian.

Considerations of table No. 3 reveal that three palynomorph genera (recovered during the present study) deviate considerably from their already established stratigraphical range of occurrence. Leptolepidites was found in the Lumshiwal Formation (Lower Cretaceous) which otherwise is only restricted to the Jurassic strata. Densipollenites has a restricted range of occurrence between Early Permian to Late Jurassic, but in the Kala Chitta Range it was identified in the Kawagarh Formation (Upper Cretaceous), whereas, Lundbladispora (purely a Triassic genus) was also isolated from the Samana Suk Limestone (Middle Jurassic). All these exceptions might be due to the reworking/stratigraphical leak or the contamination of the palynomorphs. If further studies reconfirm such instances in the same outcrops of adjoining areas, the stratigraphical occurrences of these genera (Leptolepidites Densipollenites, Lundbladispora) might be revised.

SYSTEMATIC PALYNOLOGY

Anteturma SPORITES H. Potonie, 1893 Turma TRILETES (Reinsch) Dettmann, 1963 Suprasubturma ACAVATITRILETES Dettmann, 1963 Subturma AZONOTRILETES (Luber) Dettmann, 1963 Genus *Leiotriletes* Naumova ex Isohchenko 1952 Type species *Leiotriletes sphaerotriangulatus* (Loose) Potonie and Kreump 1954 *Leiotriletes sp.* Plate No. 2b, Plate No. 5f

Description: Miospore, trilete, amb triangular or subrounded, distorted due to compression, "Y" marking not very clear, longer than $\frac{1}{2}$ radius. Exine faintly infrapunctate, up to 1 μ m.

Genus *Punctatisporites* Ibrahim 1933 Type species *Punctatisporites punctatus* (Ibrahim) Ibrahim 1933 *Punctatisporites sp.* Plate No. 4e

Description: Miospore, trilete, amb \pm circular, trilete rays relatively short, may or may not be reaching equator. Exine punctate or finely reticulate up to 1 μ m thick.

Genus *Frangospora* Venkatachala and Kar 1968 Type species *Frangospora fracta* Venkatachala and Kar *Frangospora fracta* Venkatachala and Kar Plate No. 4f

Description: Miospore, trilete, amb circular to subcircular rays $\frac{3}{4}$ radius, tapering, exine laevigate, sometimes infra-structured, unevenly thickened, outer exine splitting up by irregularly distributed cleavages giving the spore a mud crack like appearance, exine 2 µm.

Genus Foveosporites Balme 1957 Type species Foveosporites canalis Balme 1957 Foveosporites subtriangularis (Brenner) Doring Plate No. 2f

Description: Miospore, trilete, amb rounded or elongate triangular, angles broadly rounded sides straight to convex. "Y" marking extending $\frac{3}{4}$ radius. Exine infraverrucate to foveolate $\pm 3 \mu m$ thick.

Genus Verrucosisporites Ibrahim, 1933 Type species Verrucosisporites verrucosus Ibrahim Verrucosisporites narmianus Balme, 1970 Plate No. 5b

Description: Miospore, trilete, amb circular to subcircular, exine vertucose, vertucae closely set, height equal to or less than their breadth, apex truncate, average diameter of vertucae 3-4 μ m, exine up to 3 μ m thick.

Genus *Leptolepidites* Couper, 1953 Type species *Leptolepidites verrucatus* Couper *Leptolepidites eparcornatus* Plate No. 2c

Description: Miospore, trilete, amb subtriangular, sides convex to concave, biconvex in equatorial view. "Y" marking usually well developed, but not discernable in present specimen due to poor preservation. Exine very thick (up to 3 μ m). Sculptured with large irregularly shaped rounded verrucae.

Table 2 Relationship between various chemical compounds and palynomorph occurrence in the Mesozoic sucession of the Kala Chitta Range.

| Age | Formation | Sample | SiO ₂ | Fe_2O_3 | FeO | P_2O_5 | TiO ₂ | Al_2O_3 | CaO | MgO | MnO | Na ₂ O | K ₂ O | H_2O^+ | H_2O^- | CO ₂ | S | С | Total | Pal/g |
|-------------|----------------|--------|------------------|-----------|-------|----------|------------------|-----------|-------|-------|-------|-------------------|------------------|----------|----------|-----------------|-------|-------|-------|-------|
| Triassi c | Kingriali | 576 | 0.48 | 1.55 | 1.18 | - | 0.15 | 0.80 | 42.62 | 8.06 | 0.04 | 1.54 | 1.25 | 0.10 | 0.20 | 41.74 | 0.24 | 0.10 | 99.99 | 164 |
| | Samana | 703 | 3.91 | 1.13 | 0.42 | 4.41 | 3.02 | 0.47 | 43.88 | 0.20 | 0.04 | 2.45 | 1.28 | 0.10 | 0.02 | 35.59 | 0.87 | 2.21 | 100 | 1050 |
| Jurassic | Suk | 1029 | 10.92 | 1.59 | 0.16 | 0.04 | 0.30 | 0.37 | 28.04 | 16.13 | 0.04 | 1.62 | 0.68 | 0.10 | 0.12 | 37.62 | 0.06 | 2.21 | 100 | 54 |
| | | 658 | 50.70 | 1.04 | 0.12 | 0.02 | 0.30 | 6.63 | 20.50 | 2.94 | 0.07 | 1.65 | 0.68 | 0.10 | 0.09 | 14.61 | 0.15 | 0.40 | 100 | 800 |
| Jurassic / | | 669 | 16.22 | 4.82 | 0.10 | 0.80 | 0.20 | 6.98 | 35.89 | 2.42 | 0.01 | 0.21 | 0.26 | 0.10 | 0.09 | 31.73 | 0.06 | 0.10 | 99.99 | 482 |
| Creta-ceous | Chichali | 754 | 34.74 | 1.30 | 0.42 | 0.02 | 0.30 | 7.69 | 30.44 | 5.53 | 0.04 | 1.35 | 0.50 | 0.10 | 0.10 | 17.26 | 0.29 | 0.10 | 100 | 150 |
| | | 148 | 28.74 | 0.30 | 0.02 | 0.02 | 0.30 | 6.59 | 36.44 | 0.35 | 0.04 | 1.35 | 0.50 | 0.10 | 0.10 | 23.20 | 0.29 | 0.10 | 98.56 | 4080 |
| | Lumshiwal | 149 | 44.67 | 1.09 | 0.24 | 0.02 | 0.30 | 8.08 | 21.86 | 1.81 | 0.05 | 2.45 | 1.28 | 0.10 | 0.13 | 17.77 | 0.06 | 0.10 | 100 | 2400 |
| | | 187 | 15.04 | 2.79 | 0.10 | 1.91 | 0.10 | 8.43 | 38.13 | 2.36 | 0.02 | 0.86 | 0.23 | 1.51 | 0.62 | 26.26 | 0.15 | 1.30 | 99.99 | 1000 |
| | | 188 | 47.40 | 8.16 | 2.42 | 0.05 | 1.50 | 8.99 | 12.33 | 1.61 | 0.04 | 1.53 | 1.96 | 0.10 | 3.30 | 10.30 | 0.19 | 0.10 | 99.98 | 9000 |
| Lower | | 209 | 20.42 | 5.71 | 0.44 | 0.09 | 0.42 | 9.83 | 33.94 | 5.91 | 0.09 | 1.53 | 3.01 | 0.10 | 2.00 | 16.35 | 0.06 | 0.10 | 100 | 6200 |
| | | 210 | 14.61 | 9.00 | 2.78 | 0.17 | 0.97 | 12.58 | 39.11 | 3.22 | 0.01 | 3.45 | 4.20 | 1.68 | 1.29 | 6.17 | 0.03 | 0.20 | 99.52 | 10100 |
| | I | 661 | 45.19 | 5.15 | 0.32 | 0.09 | 0.20 | 1.33 | 21.16 | 1.35 | 0.06 | 2.22 | 2.41 | 0.10 | 3.30 | 16.90 | 0.04 | 0.20 | 100 | 8000 |
| | I | 665 | 29.00 | 1.30 | 0.32 | 0.02 | 0.30 | 5.69 | 28.00 | 5.60 | 0.04 | 1.32 | 0.06 | 0.10 | 0.20 | 27.10 | 0.26 | 0.10 | 99.94 | 2000 |
| | I | 666 | 20.20 | 1.23 | 0.32 | 0.09 | 0.08 | 5.20 | 33.25 | 2.31 | 0.06 | 1.35 | 0.68 | 0.10 | 0.08 | 33.31 | 0.06 | 0.10 | 98.42 | 4200 |
| Creta-ceous | | 756 | 9.33 | 0.09 | 1.80 | 0.11 | 0.71 | 10.15 | 48.82 | 6.15 | 0.05 | 3.36 | 3.89 | 0.93 | 2.92 | 6.62 | 0.20 | 2.21 | 99.34 | 7100 |
| | | 170 | 24.70 | 1.30 | 0.42 | 0.01 | 0.20 | 2.69 | 38.45 | 6.35 | 0.04 | 2.17 | 2.56 | 0.10 | 0.32 | 19.30 | 0.07 | 1.30 | 99.98 | 400 |
| | I | 171 | 16.74 | 1.24 | 0.10 | 0.02 | 0.38 | 1.24 | 40.72 | 2.97 | 0.06 | 2.22 | 2.65 | 0.10 | 0.14 | 30.41 | 0.87 | 1.30 | 100 | 192 |
| Upper | I | 1236 | 20.15 | 3.25 | 2.17 | 0.05 | 0.20 | 4.30 | 35.77 | 1.64 | 0.02 | 2.11 | 2.71 | 0.10 | 0.11 | 27.02 | 0.16 | 0.20 | 99.96 | 1000 |
| Creta-ceous | Kawagarh | 1237 | 18.20 | 2.20 | 0.17 | 0.04 | 0.20 | 4.60 | 38.40 | 2.02 | 0.04 | 2.05 | 1.51 | 0.10 | 0.13 | 30.08 | 0.06 | 0.20 | 100 | 984 |
| | r-values | | +0.39 | +0.66 | +0.60 | -0.22 | +0.24 | +0.56 | -0.02 | -0.23 | -0.20 | +0.50 | +0.61 | +0.11 | +0.83 | -0.70 | -0.30 | -0.20 | | |
| (Co-effici | ient of correl | ation) | ns | ** | ** | ns | ns | ** | ns | ns | ns | * | ** | ns | ** | ** | ns | ns | - | - |

Legend: ns= not significant at $\alpha 0.05$ Pal/g = Palynomorph per gram of sample *= significant at $\alpha 0.05$

**= highly significant at α 0.05 (significant at α 0.01)

| | AGE | Т | RIASSIC | | J | URASSIC | | CRETACEOUS | | | | |
|-------|----------------------|-------|---------|------|-------|---------|------|------------|--------|----------|--|--|
| S. No | Genus | Early | Middle | Late | Early | Middle | Late | Early | Middle | Late | | |
| 1 | Affropollis | | | | | | | | | | | |
| 2 | Densoisporites | | | | | | | | | | | |
| 3 | Alisporites | « | | | | | | | | ~ | | |
| 4 | Frangospora | | | | | | | | | | | |
| 5 | Inaperturopollenites | | | | | | | | - | * | | |
| 6 | Leptolepidites | | | | | | | | | | | |
| 7 | Leiotriletes | « | | | | | | | | » | | |
| 8 | Laevigatosporites | | | | | | | | | | | |
| 9 | Minutosaccus | - | | | | 4.) | | | | | | |
| 10 | Ovalipollis | | | 1.2 | | | | | | | | |
| 11 | Podocarpites | | | | | | | | | | | |
| 12 | Punctatisporites | | | | | | | | | | | |
| 13 | Succinctisporites | | | | | | | | | | | |
| 14 | Valiasaccites | | | | | | | | | | | |
| 15 | Foveosporites | | | | | | | | | | | |
| 16 | Crybelosporites | | | | | | | | | | | |
| 17 | Verrucosisporites | | | | | | | | | | | |
| 18 | Cyclobaculisporites | | | | | | | | | | | |
| 19 | Inferopollenites | | | | | | | | | | | |
| 20 | Densipollenites | | | | | | | | | | | |
| 21 | Pinuspollenites | | | | | | | | | | | |
| 22 | Aequitriradites | | | | | | | | | | | |
| 23 | Lundbladispora | | | | | | | | | | | |
| 24 | Pterospermella | | | | | | | | | | | |
| 25 | Goubinispora | | | | | ч. | | | | | | |

 Table 3

 Tentative range of occurrence of some stratigraphically important palynomorph genera recovered from the Mesozoic Rocks of the Kala Chitta Range.

Age not clearly established.

≪ Extending into older rocks

>> Extending into younger rocks

Subinfraturma NODATI Dybova and Jackovicz (1957) Genus Cyclobaculisporites Bharadwaj 1955

Type species *Cyclobaculisporites* grandiverrucosus (Kosanke) Bharadwaj

Cyclobaculisporites minutus Kar 1968 Plate No. 5a

Description: Miospore, trilete, amb circular to subcircular, trilete rays indistinct, less than _ radius, exine surface thickly set with baculae of uniform length but widely ranging width, interbacular spaces very narrow, discernable only under oil immersion objective.

Infraturma CINGULATI (Potonie and Klaus) Dettmann, 1963

Genus *Densoisporites* Weyland and Krieger, 1953 Type species *Densoisporites velatus* Weyland and Krieger

Densoisporites nejburgii (Schulze) Balme Plate No. 2a

Description: Miospore, trilete, amb rounded, triangular, "Y" mark distinct, rays extending to or into the flange, with narrow raised lips fading towards equator. Exine up to 2 μ m thick, with surface finely granulate to chagrenate. The exoexine is proximally and distally \pm tightly fitting, but equatorially cavate and sculptured by numerous short plications and wrinkles.

Genus *Crybelosporites* Dettmann 1963 Type species *Crybelosporites striatus* (Cookson and Dettmann) Dettmann *Crybelosporites pannuceus* (Brenner) Srivestava Plate No. 3e

Description: Miospore, trilete, amb spheroidal to ellipsoidal, exine stratified consisting of a smooth homogenous inner layer enclosed within a two layered, proximally cavate, structured 'sculptine', outer layer of sculptine without a trilete aperture, proximally detached from trilete, inner layers; and with a reticulate, rugulate, or foveolate (OL) surface pattern.

> Genus Lundbladispora Balme 1963 Type species Lundbladispora willmotti Balme Lundbladispora willmotti Balme

Description: Miospore, trilete, amb subcircular to subtriangular, exine cavate, a finely structured exoexine enclosing a thin walled intexine, "Y" marking usually well developed.

> Turma MONOLETES Ibrahim 1933 Suprasubturma ACAVATOMONOLETES Dettmann, 1963 Subturma AZONOMONOLETES Luber, 1935

Genus *Laevigatosporites* Ibrahim, 1933 Type species *Laevigatosporites vulgaris* Ibrahim *Laevigatosporites sp.* Plate No. 2j, k

Description: Miospore, monolete, amb broadly oval, tetrad scar distinct, alete extending $\frac{3}{4}$ length of spore, commissure thin, bordered by distinct labra. Exine up to 3 μ m thick, laevigate to chagrenate.

Turma ALETES Ibrahim Genus Inaperturopollenites Pflug in Thomson and Pflug 1953 Type species Inaperturopollenites dubius (Potonie and Venkatachala) Thomson and Pflug 1953 Inaperturopollenites sp. Plate No. 3b

Description: Pollen grain, germinal apparatus or suture indistinct, amb circular or subcircular. Exine thin up to 1 μ m thick, infrapunctate with many secondary folds.

Anteturma POLLENITES Potonie 1931 Turma SACCITES Erdtman 1947 Subturma MONOSACCITES (Chitaley) Potonie and Kremp 1954 Genus *Enzonalasporites* Leschik 1956a Type Species *Enzonalasporites vigens* Leschik *Enzonalasporites vigens* Leschik Plate No. 4d

Description:- Pollen grain, monosaccate, overall amb circular, corpus circular to sub circular not very well demarcated due to the poor state of preservation in the given specimen, exine of corpus rugulose, alete, velum not very sharply differentiated from central body.

> Genus Goubinispora Type species Goubinispora Goubinispora sp.

Description: Pollen grain, monosaccate, central body diffuse, not sharply delineated, alete, saccus slightly lobed, exoexine of saccus infrareticulate, up to 2µm thick.

> Subturma DISACCITES Cooksoon, 1947 Infraturma STRIATITI Pant, 1954 Genus Infernopollenites Scheuring 1970 Type species Infernopollenites sulcatus Scheuring Infernopollenites sp. Plate No. 5h

Description:- Pollen grain, bisaccate, diploxylonoid, central body circular to subcircular, usually three broad longitudinal taeniae that usually are closely joined or are separated by narrow ectexinal fissures covering the proximal hemisphere are characteristic for this genus, but these are not sharply delineated here due to bad state of preservation. Distal bases of sacci are straight. Exine of corpus infrapunctate up to 1.5 μ m thick. Exoexine of sacci 2 μ m.

Non Taeniate Bisaccate

Genus *Pinuspollenites* Raatz 1938, Potonie 1958 Type species *Pinuspollenites labdacus* (Potonie) Raatz ex Potonie 1958 *Pinuspollenites sp.* Plate No. 5c

Description:- Pollen grain, bisaccate, slightly diploxylonoid, corpus circular to subcircular, without germinal suture, exine of corpus laevigate up to 1.5 μ m thick. Sacci small, attachment ventral, exoexine of sacci 2 μ m thick.

Genus Valiasaccites Bose and Kar 1966 Type species Valiasaccites validus Bose and Kar Valiasaccites validus Bose & Kar Plate No. 2i

Description:-Pollen grain, bisaccate. haploxylonoid, overall amb ± oval, central body oval, usually with two longitudinal lateral ridges which are not detectable in the present specimen. Exine of corpus infravermiculate to infrareticulate. Monolete mark on corpus absent. Proximal saccus attachment zone along equator, distally subequatorial and associated with a semilunar Sacci body. semicircular, coarsely intrareticulate (makes up to 2 µm), attachment of sacci lateral with slight ventral displacement. Exoexine of sacci up to 2 µm thick.

> Genus *Podocarpites* Bolkhovitina 1956 Type species *Podocarpites flacciformis* (Malj.) ex Bolkhovitina *Podocarpites sp.* Plate No. 1h

Description:- Pollen grain, bisaccate haploxylonoid to slightly diploxylonoid. Central body circular to subcircular, alete. Exine of corpus infrapunctate up to 1.5 μ m thick. Attachment of sacci lateral with slight ventral displacement. Exoexine of sacci up to 2 μ m thick.

Genus *Ovalipollis* Krutrch 1955 Type species *Ovalipollis ovalis* Krutrch *Ovalipolis ovalis* Krutrch Plate No. 3c

Description:- Pollen grain, bisaccate, overall amb slender or broadly oval, in part rhombic, usually a major fold runs longitudinally on one face from one end to other, the pointed ends of oval central body are capped by weakly developed sacci that possess a pronounced rodlet structure and have a smooth surface. Exine of central body up to 1.5 μ m thick whereas that of sacci up to 2 μ m thick.

> Genus *Minutosaccus* Madler 1964 Type species *Minutosaccus acutus* Madler 1984 *Minutosaccus crenulatus* Dobly Plate No. 2g

Description:- Pollen grain, bisaccate, haploxylonoid, overall amb broadly fusiform, central body circular to subcircular, strongly cutinized and small, sacci distally displaced, usually with sharply delineated sulcus which is detectable in actual specimen under oil immersion objective upon careful L-O analysis but not in the photomicrograph provided. Exine of corpus laevigate up to 2 µm thick, exoexine of sacci up to 1.5 µm thick.

> Genus Alisporites Daugherty emend Nilsson, 1958 Type Species Alisporites opii Daugherty Alisporites grandis (Cookson) Dettman Plate No. 1g

Description:- Pollen grain, bisaccate, haploxylonoid, central body circular to subcircular. Exine of corpus 1 μ m thick laevigate. Sacci attached laterally with slight distal inclination. Exoexine of sacci infra reticulate up to 1.5 μ m thick.

Non-Vesiculate Tectate Pollen

Genus Afropollis Doyle, Jardine and Doerenkamp 1982

Type species *Afropollis Jardina* (Brenner) Doyle *Afropollis jardina* (Brenner) Doyle Plate No. 3f

Description:- Pollen grain, overall amb spheroidal, radially symmetrical, coarsely reticulate to ruguloreticulate, semitectate. Due to corrosion and poor preservation, differentiation of tectum is not so sharp in the photomicrograph provided. It was however, discernable under oil immersion objective upon careful L-O analysis in the actual specimen. Lumina on most of the grain surface 2-5 μ m maximum dimension.

GEOCHEMISTRY AND PALYNOMORPH PRODUCTIVITY

During the present study an attempt has been made to investigate whether or not the chemical composition (or any change in it) of the parent sediment affects productivity and preservation of palynomorphs in the Mesozoic rocks of the Kala Chitta Range. However, if a particular sediment is not productive (containing no palynomorphs) no method is devised to date to judge whether it originally contained palynomorphs or not, or the microfossils disappeared later due to some postdepostitional hazard(s). The author has personally observed that if a zone of a particular formation is highly productive at one level, it may become poorly productive or totally barren at another horizon. As indicated by several workers it may be due to the changes in depositional environments (Darrell and Hart, 1979), thermal effects (Dow, 1977) or other post depositional hazards (Bonny, 1976, 1978).

Some other sedimentological factors in combination with the associated vegetational pattern on nearby land and the differential rate of transportation of various palynomorphs to the depositional site may affect occurrence and preservation of palynoflora in rocks (Smith, 1962; Chaloner and Muir, 1968; Phillips et al., 1974; Scott and King, 1981). But the possibility of the fact that any change(s) in the chemical composition of the parent sediment may affect palynomorph productivity cannot be ruled out totally.

Twenty rock samples viz KQ 575 (Kingriali Formation); KQ 703, KQ 1029 (Samana Suk Limestone); KQ 658, KQ 669, KQ 754 (Chichali Formation); KQ 148, KQ 149, KQ 187, KQ 188, KQ 209, KQ 210, KQ 661, KQ 665, KQ 666, KQ 756 (Lumshiwal Formation); KQ 1236, KQ 1237, KQ 170, KQ 171 (Kawagarh Formation) were randomly selected to study such phenomenon.

Chemical analysis

The contents of SiO_2 , H_2O+ , H_2O- were determined gravimetrically. Al_2O_3 , MnO, CaO and MgO were determined by atomic absorption and spectrophotometry.

The amounts of TiO_2 and P_2O_5 were determined by spectrophotometry. FeO and inorganic carbon were determined volumetrically. The alkalies and organic carbon were determined by C and S analyser.

Statistical analysis

Coefficient of correlation (r-values) were calculated between percentages of an elemental oxide and number of palynomorphs/g in different samples to check whether any relationship exists between the two independent variables viz: percentage of an elemental oxide and number of polynomorphs per gram of samples (productivity).

Table No. 2 shows the sample wise percentage of elemental oxides and number of palynomorphs/g of different samples and the r-values.

It has been observed that FeO, Fe_2O_3 , Al_2O_3 , K_2O show highly significant positive correlation between the two independent variables indicating a highly

favourable chemical environment for palynomorph preservability. Na₂O also shows significant positive correlation. Negative correlation is also observed in P₂O₅, CaO, MgO, MnO, though not significant at 5% level of significance. However, it may indicate that oxidizing environments such as that of P₂O₅ are highly unfavourable for palynomorph preservation. Similarly, CaO and MgO, which are components of dolomites, are also unsuitable for palynomorph preservability. Similar observations have also been made by Masood et al. (1995) on the samples from the Amb Formation (Permian) of the Salt Range.

The above study indicates that the higher amounts of FeO, Fe_2O_3 , Al_2O_3 , K_2O , Na_2O , SiO_2 TiO₂ favour higher palynomorph occurrence whereas P_2O_5 , CaO, MgO and MnO₂ impart low productivity.

PALEOCLIMATOLOGY AND DEPOSITIONAL ENVIRONMENTS

The climate experienced by an area at a particular time depends upon its continentality and latitudinal position. Fossil spores and pollen exhibit various morphographic characters with specific configurations and structures, each with a precise function. These morphographic features transcend taxonomic delimitations of the dispersed spores and pollen. The behaviour pattern of significant morphographic events are evaluated for their sensitivity to climate. The combination of characters in a given palynoflora is diagnostic for that period and it has been considered for climatic interpretation. Since the palynoflora recovered during the present investigation was not well preserved, it was not possible in most cases to interpret morphographic features associated with palynomorphs correctly, and hence accurate climatic interpretation at various stratigraphic levels was not possible.

Considering Bharadwaj scheme (1966) it is possible to some extent to comment on vegetational history and then to elucidate the depositional environment on a broad scale. According to Bharadwaj scheme -trilete and monolete miospores represent Cryptogams striated bisaccates, Glossopteroids; non-striated bisaccates, Conifers, trilete monosaccates, Gangamopteroids, monosulcates, Cycadoginkops and alete monosaccates represent Cordaitales.

Absence of monosaccate in all samples except in the Kingriali Formation (i.e. *Goubinispora*) is noteworthy (Table No. 1), indicating that none of the members of *Pteridospermaeles* existed within the close vicinity of the depositional site from Jurassic to Upper Cretaceous periods.

CONCLUSION

Assemblages from Kingriali Formation, the Samana Suk Limestone and the Chichali Formations are characterised by poor preservation, lack of diversity and the presence of enormous small spinose acritarch and indistinctive leiospheres. The Kingriali Formation also species of Lundbladispora contains few and Goubinispora. Both these types indicate tropical to subtropical hot climate with medium to high humidity in closely located land area, whereas, the palynoflora of the Samana Suk Limestone and the Chichali Formation include great abundance of cryptogams, which must be present very near to shore in an upland area with typical temperate to subtemperate (low cooling) with medium to high humidity.

Assemblages of the Lumshiwal Formation are much more diverse than those recovered from any other formation. The distributional heterogeneity is evident from consideration of Table 1. Several pollen species were found only in the Lumshiwal Formation and the diversification of spore and pollen assemblage is due mainly to an increase in the variety of trilete and bisaccate palynomorphs.

Taken in its entirety a strongly regressive phase, which may of course have been associated with wide climatic changes, have been observed. Regression would be expected to manifest itself by an increase in the number of plant microfossils reaching the depositional area. A decrease in the number and veriety of the acritarchs and the diversification of the assemblages associated with large number of pollen and spores derived from subordinate or locally restricted floral elements are the characteristics of the assemblages from the Lumshiwal Formation.

There is a marked decline in this trend in the Kawagarh Formation. However, the Kawagarh Formation is palynologically more rich as compared to the Kingriali Formation or the Samana Suk Limestone.

REFERENCES

- Bharadwaj, D.C. 1966. Distribution of Spores and Pollen grains dispersed in the Lower Gondwana Formations of India. Symp. Florist. Strat. Gondwanaland. Spec. Session Publ: 69-84.
- Blome, C.R. and Albert, N.R. 1985. Carbonate Concentrations: an ideal Sedimentary Host for Microfossils. *Geology* **13**: 212-215.
- Bonny, A.P. 1976. Recruitment of pollen to the sesten and sediment of some English Lade Distrct lakes: *Jour. Ecol.* **64**: 859-87.
- Butt, A.A., Qureshi, M.K.A., Masood, R. and Ghazi, S. 2008. Palynology of the Mesozoic Succession of the Kala Chitta Range, Pakistan. Abstract Volume 33rd Internat. Geol. Congr. Oslo, Norway.
- Chaloner, W.G., Muir, M. 1968. Spores and Floras in coal and coal bearing strata, D.G. Murchinson & T.S. Westall (eds.), 127-146 Oliver & Boyd Edinburgh.
- Cotter, G. De. P. 1933. The Geology of the part of the Attock District, West of Long. 72°45′ E. Mem. Geol. Surv. India 55: 63-161.
- Darrell, J.H. and G.F. Hart. 1979. Environmental determinations using absolute miospore frequency, Mississippi River delta. Geol. Soc. Amer. Bull. 81: 2513-18.
- Dow, W.G. 1977. Kerogen studies and geological interpretations. Jour. Geochem. Explor. 7: 79-99.
- Fatmi, A.N., 1973. Lithostratigraphic units of Kohat-Potwar Province, Indus Basin, Pakistan, *Mem.Geol. Surv: Pakistan* **10**:1-80.
- Jansonius, J. and Hills, H. 1976. Fossil genera file of Palynomorphs, Univ. Calgary, Spec. Publ.
- Masood, K.R., Qureshi, K.A., Hussain, Z. and Parveen, M. 1995. Palynomorph occurrence in relation to geochemistry, in the Amb Formation (Artiiskian), Zaluch Gorge, Salt Range, Pakistan. *Pak. Jour. Hydrocarbon Research* 7(1): 61-70.
- Phillips, T.L., R.A. Peppers, M.J. Avcin and P.F. Langhnan. 1974. Fossil plants and coal: patterns of chang in Pennsylvanian coal swamps of Illinois Basin. *Science* **187**(4144), 1367-1369.
- Scott, A.C., J. Galtier and G. Clayton. 1985. A new late Tournaisian (lower carboniferous) Flora from the Kilpatrich Hills, Scotland. Rev. Paleobot. Palynol.: 44: 81-89.
- Smith, A.H.V. 1962. The palaeoecology of Carboniferous peats based on miospores and petrography of bituminous coals. *Proc. Yorkshir Geol. Soc.* 33: 423-474.